

Modelling Aquatic Ecosystems

Course 701-0426-00

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1. Introduction, principles of modelling environmental systems, mass balance in a mixed reactor, process table notation, simple lake plankton model
Exercise: R, ecosim-package, simple lake plankton model
Exercise: lake phytoplankton-zooplankton model
2. Process stoichiometry Exercises: analytical solution, calculation with stoichcalc
3. Biological processes in lakes
4. Physical processes in lakes, mass balance in multi-box and continuous systems Exercise: structured, biogeochemical-ecological lake model
Assignments: build your own model by implementing model extensions
5. Physical processes in rivers, bacterial growth, river model for benthic populations Exercise: river model for benthic populations, nutrients and oxygen
6. Stochasticity, uncertainty, Parameter estimation
Exercise: uncertainty, stochasticity
7. Existing models and applications in research and practice, examples and case studies, preparation of the oral exam, feedback

Today:

- learn how to formulate mass balances in multi-box (**chapter 3.3**) (and continuous, **3.4**) systems
- get a basic understanding of mixing and transport processes in lakes (**chapter 6.1.1**)
- introduction to the biogeochemical – ecological lake model (**11.4**)

Assignments of topics for own model

There will be an **oral exam in the second week** after the semester
11./12. June 25 (10. or 4./5. June if needed)

It is your responsibility to register/deregister in time.

During the semester you will develop and implement your own model
(in teams of two people), interpret simulation results and perform a sensitivity
analysis. We will assign topics today.

Deadline for initial code submission: 08.05.25

Deadline for submission of R-files, results and interpretation: 23.05.25

This is **mandatory for being admitted** to the exam!

In the oral exam we will start with questions about your model before moving
on to other topics.

Please use the time in the exercises to ask questions and get help!
Don't do it last minute.

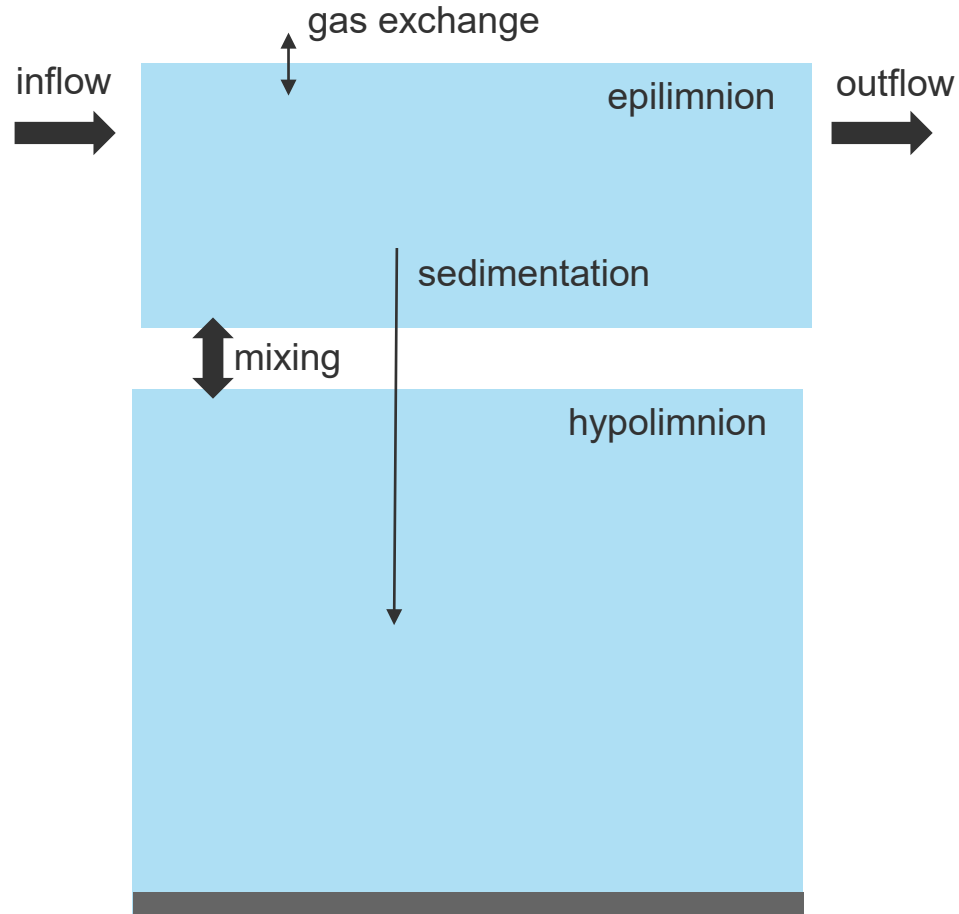
State variables:

ALG, ZOO,

HPO_4^{2-} , NH_4^+ , NO_3^- , O_2 ,

POM: degradable/inert,

suspended/sedimented



(topic of today and next exercise)

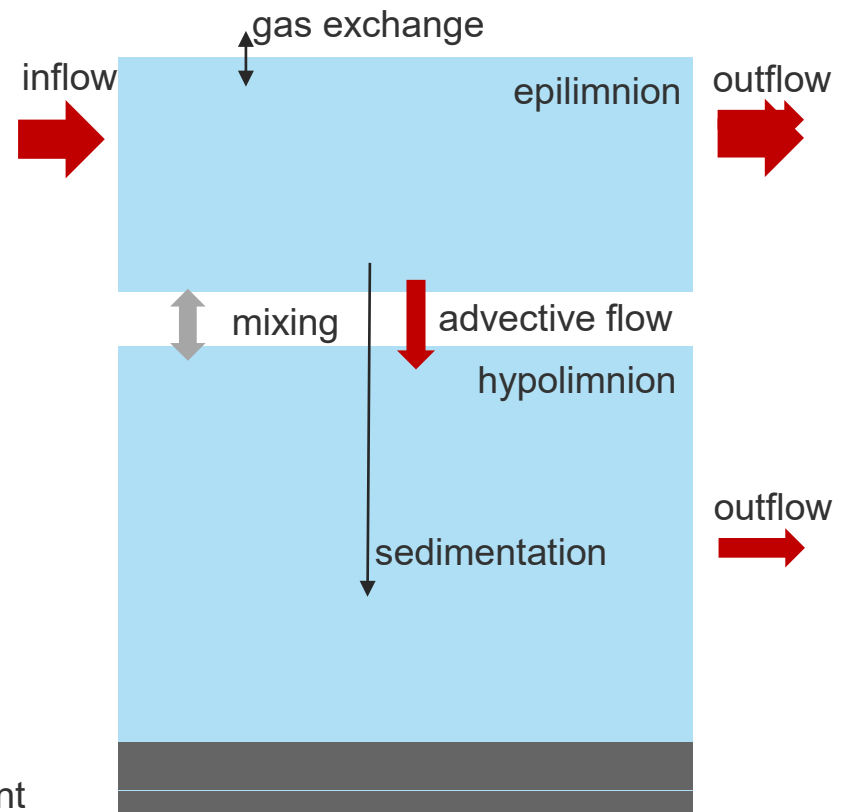
Topic 1: scenarios for phosphorus reduction in a lake: extraction of deep water and reduction of inflow

Motivation

Assess management scenarios to reduce eutrophication of a lake

Assignment

- Start from model 11.4
- Add a deep water extraction (outflow from hypolimnion)
- Investigate the following scenarios separately:
 1. Reduction of the inflow concentration of phosphorus by 50%
 2. deep water extraction: 50% of the inflow leaves the lake from hypolimnion (volume stays constant)
- Analyse the effect of both measures with a small or a large deposit of POMD in the sediment



Goal: Compare the effectiveness of different management scenarios for different sediment conditions and estimate how long it will take to see the effects.

Topic 2: insecticide contamination and coexistence of two zooplankton groups

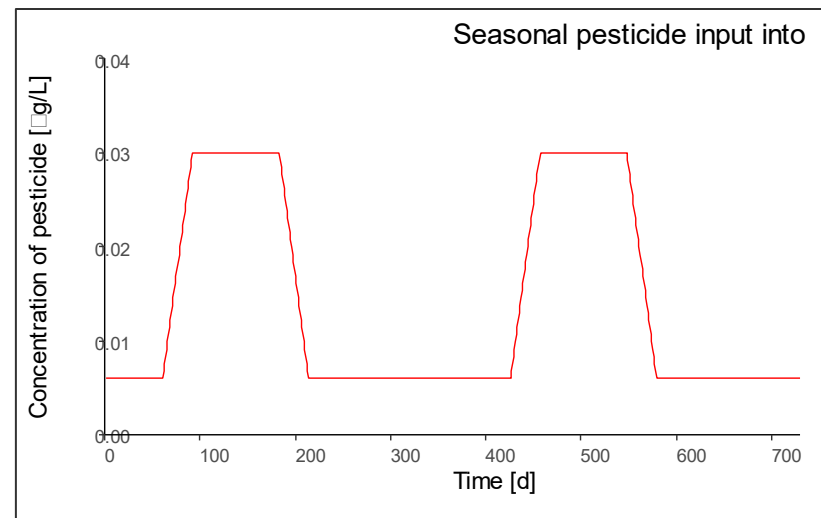
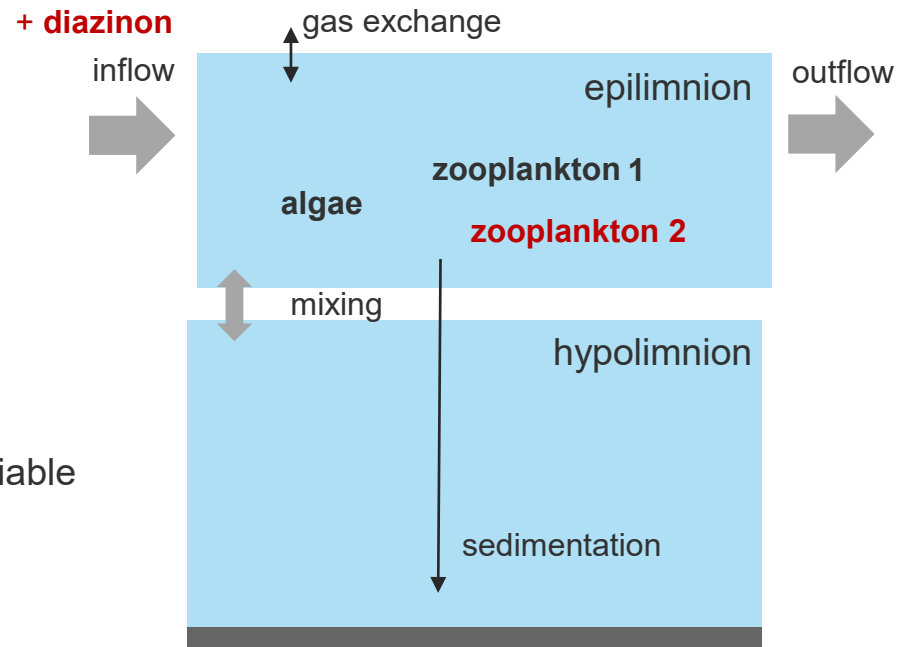
Motivation

Simulate effects of micropollutants
and analyze coexistence of multiple groups

Assignment

- start from model 11.4
- introduce an insecticide (diazinon) as state variable
- introduce two zooplankton groups differing in
 - half-saturation constant for food limitation
 - sensitivity towards insecticide

Goal: Try to achieve long-term coexistence
between the two zooplankton groups
by varying the parameters



Motivation

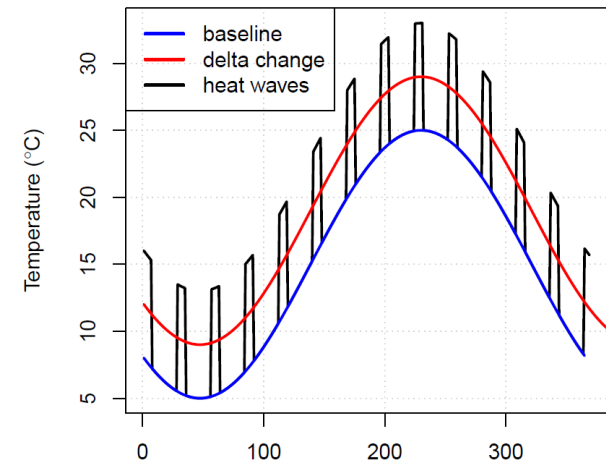
Simulate effects of temperature change scenarios.

Assignment

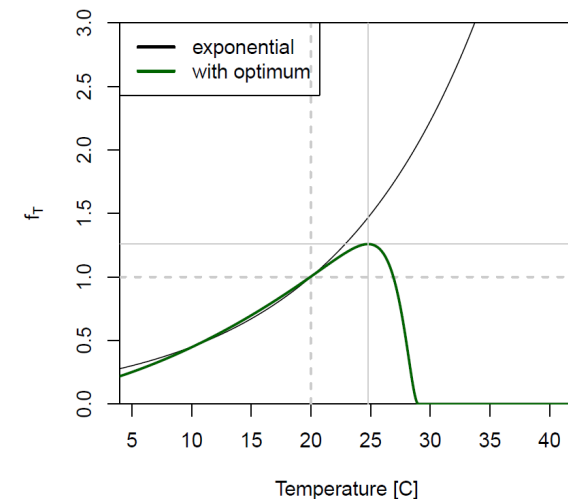
- start from model 11.4
- implement three different temperature scenarios
- change the temperature dependence of the growth rate of zoo plankton

Goal: Analyze the effects of the scenarios on zooplankton and the other state variables and reflect about how realistic the results are.

Temperature scenarios



temperature dependence functions



What do you already know about transport processes in aquatic systems?

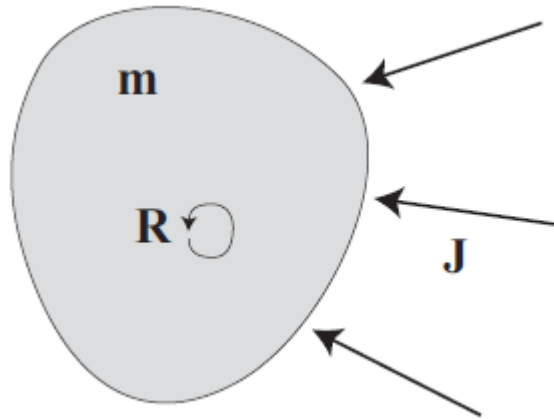


Transport processes

- Advection:** Transport of a substance by directed bulk motion of a fluid, in our case **transport with water flow**.
- Sedimentation:** Directed transport of particles in a fluid (water) caused by **gravitational forces** rather than by water flow.
- Diffusion:** Undirected transport (mixing) due to molecular motion (**molecular diffusion**) or turbulent eddies (**turbulent diffusion**).
- Dispersion:** Mixing in flow direction due to transverse diffusion of particles between flow paths with different advective velocity (“shear flow dispersion”, see transport and mixing in rivers).

Note that due to the large range of eddy sizes in turbulent flow, the distinction of “directed” (advection) and “undirected” flow may depend on the scale of a substance patch (see example on horizontal lake mixing)

General Principles and One Box Model



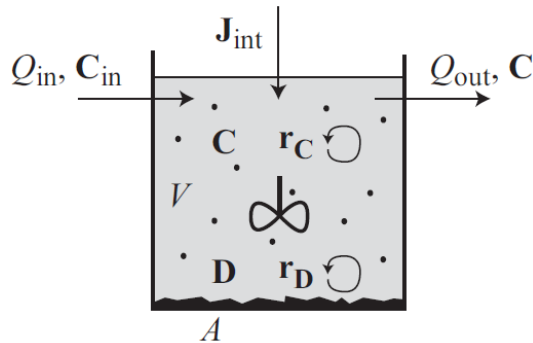
integral form:

$$\mathbf{m}(t_{\text{end}}) = \mathbf{m}(t_{\text{ini}}) + \int_{t_{\text{ini}}}^{t_{\text{end}}} \mathbf{J}(t) dt + \int_{t_{\text{ini}}}^{t_{\text{end}}} \mathbf{R}(t) dt$$

differential form:

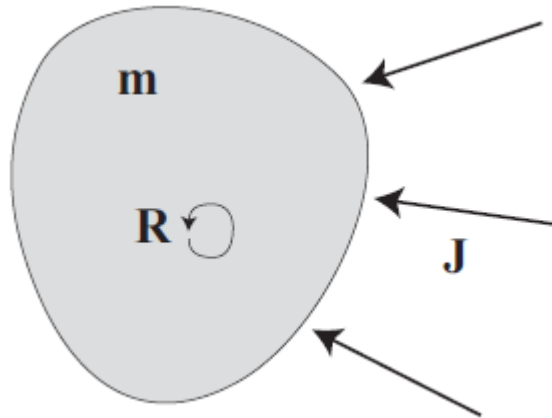
$$\frac{d\mathbf{m}}{dt}(t) = \mathbf{J}(t) + \mathbf{R}(t)$$

special case: one box



$$\mathbf{m} = \begin{pmatrix} V \\ VC_1 \\ VC_2 \\ \vdots \\ VC_{n_v} \\ AD_1 \\ AD_2 \\ \vdots \\ AD_{n_a} \end{pmatrix}, \quad \mathbf{J} = \begin{pmatrix} Q_{\text{in}} - Q_{\text{out}} \\ Q_{\text{in}}C_{\text{in},1} - Q_{\text{out}}C_1 + J_{\text{int},1} \\ Q_{\text{in}}C_{\text{in},2} - Q_{\text{out}}C_2 + J_{\text{int},2} \\ \vdots \\ Q_{\text{in}}C_{\text{in},n_v} - Q_{\text{out}}C_{n_v} + J_{\text{int},n_v} \\ 0 \\ 0 \\ \vdots \\ 0 \end{pmatrix}, \quad \mathbf{R} = \begin{pmatrix} 0 \\ Vr_{C1} \\ Vr_{C2} \\ \vdots \\ Vr_{C_{n_v}} \\ Ar_{D1} \\ Ar_{D2} \\ \vdots \\ Ar_{D_{n_a}} \end{pmatrix}$$

General Principles and One Box Model



integral form:

$$m(t_{\text{end}}) = m(t_{\text{ini}}) + \int_{t_{\text{ini}}}^{t_{\text{end}}} \mathbf{J}(t) dt + \int_{t_{\text{ini}}}^{t_{\text{end}}} \mathbf{R}(t) dt$$

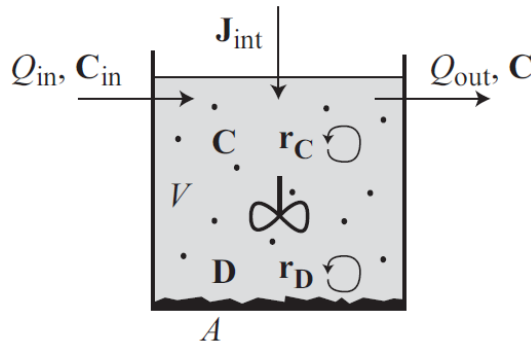
differential form:

$$\frac{dm}{dt}(t) = \mathbf{J}(t) + \mathbf{R}(t)$$

volume change as difference
of inflow - outflow

mass inflow – outflow
caused by water flow
(dissolved/suspended substances)

special case: one box



$$\frac{dV}{dt} = Q_{\text{in}} - Q_{\text{out}}$$

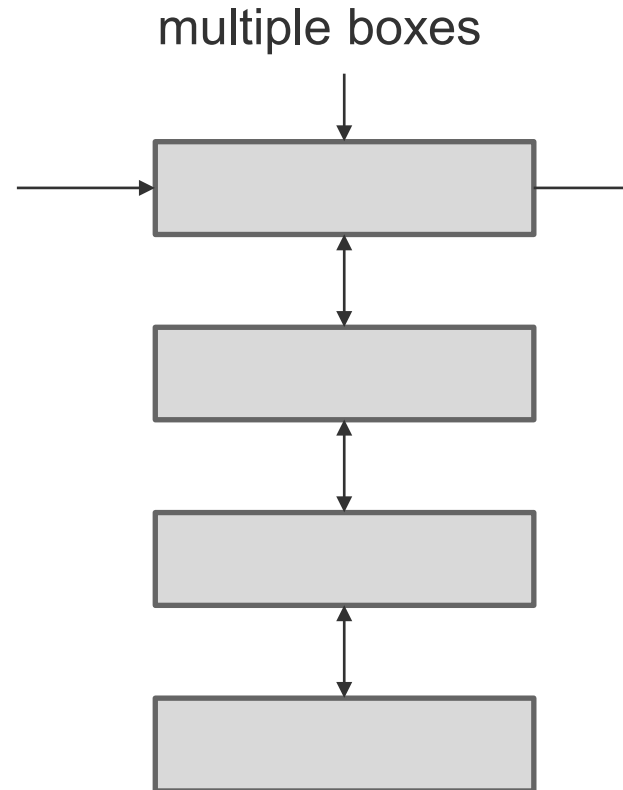
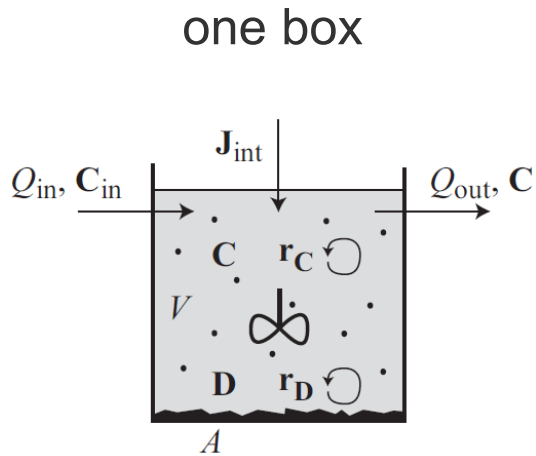
mass flow across interfaces
not related to water flow (e.g.
sedimentation, gas exchange)

$$\frac{d}{dt}(VC) = Q_{\text{in}}C_{\text{in}} - Q_{\text{out}}C + J_{\text{int}} + Vr_C$$

$$\frac{d}{dt}(AD) = Ar_D$$

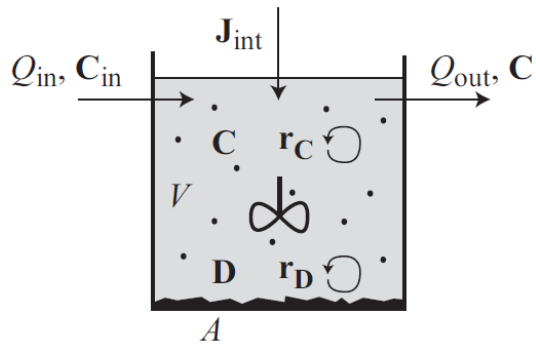
effect of
transformation processes

Multi-Reactor System

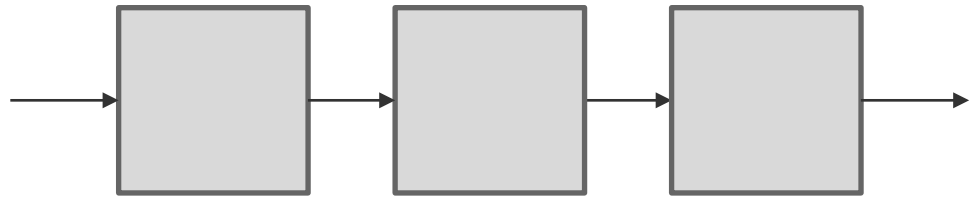


Multi-Reactor System

one box

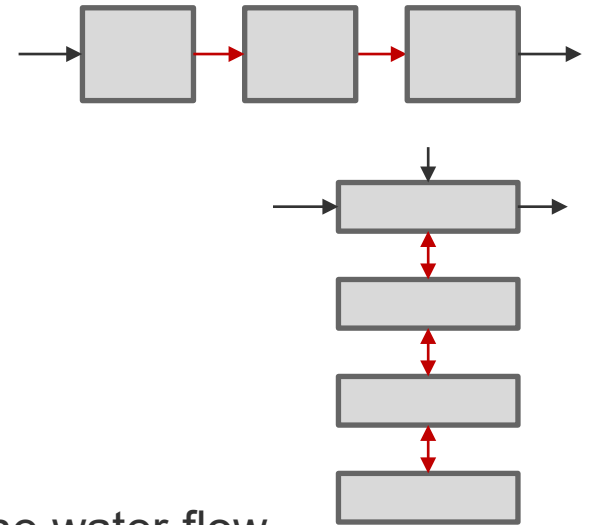


multiple boxes



New element: **links** between the reactors →

- need a definition of the logical flow direction
(**from** box x **to** box y)
- the real flow direction depends if $Q_{\text{link}} > 0$ or < 0
- can include **advective transport** of substances with the water flow or **directed** transport such as sedimentation, where the flow depends only on the **concentration in the originating box**
- and **diffusive transport**, independent from the water flow, where the flow depends on **concentration differences between the two boxes**



directed mass flows

Transport proportional to the concentration of the originating reactor:

$$J_{l,j}^{\text{adv}} = q_{l,j}^{\text{adv}} \cdot C_j^{\text{if}(q_{l,j}^{\text{adv}}, k_l^{\text{from}}, k_l^{\text{to}})}$$

$$q_{l,j}^{\text{adv}} = A v_{\text{sed},j}$$

flow direction logical:
notation to choose concentration
from originating reactor

$$\text{if}(Q, k_1, k_2) = \begin{cases} k_1 & \text{if } Q \geq 0 \\ k_2 & \text{if } Q < 0 \end{cases}$$

sedimentation:
proportionality coefficient [L³/T]
equal to cross-sectional area A
times sedimentation velocity v_{sed}

diffusive mass flows

Diffusive exchange leads to a flux proportional to the concentration difference across the interface:

$$J_{l,j}^{\text{diff}} = q_{l,j}^{\text{diff}} \cdot (C_j^{k_{\text{from}}} - C_j^{k_{\text{to}}}) \quad q_{l,j}^{\text{diff}} = A \frac{D_j}{L_{\text{diff}}}$$

q^{diff} : proportionality coefficient [L^3/T]
equal to the cross-sectional area A
times the diffusion coefficient D
divided by the thickness of the diffusive
layer L_{diff}

Mass Balance Equations Multiple Reactors

water flow from reactor
 k into other reactors

water flow from other
reactors into reactor k

$$\frac{dV^k}{dt} = \boxed{Q_{\text{in}}^k - Q_{\text{out}}^k} - \sum_{l \text{ with } k_l^{\text{from}}=k} Q_l + \sum_{l \text{ with } k_l^{\text{to}}=k} Q_l$$

$$\frac{d}{dt} (V^k \mathbf{C}^k) = \boxed{Q_{\text{in}}^k \mathbf{C}_{\text{in}}^k - Q_{\text{out}}^k \mathbf{C}^k + \mathbf{J}_{\text{int}}^k + V^k \mathbf{r}_{\mathbf{C}}^k}$$

equations for each
reactor k as in the
1 box case

$$\begin{aligned} & - \sum_{l \text{ with } k_l^{\text{from}}=k} \left(Q_l \mathbf{C}^{\text{if}(Q_l, k_l^{\text{from}}, k_l^{\text{to}})} + \mathbf{J}_l \right) \\ & + \sum_{l \text{ with } k_l^{\text{to}}=k} \left(Q_l \mathbf{C}^{\text{if}(Q_l, k_l^{\text{from}}, k_l^{\text{to}})} + \mathbf{J}_l \right) \end{aligned}$$

mass flow from
reactor k into
other reactors
mass flow from
other reactors
into reactor k

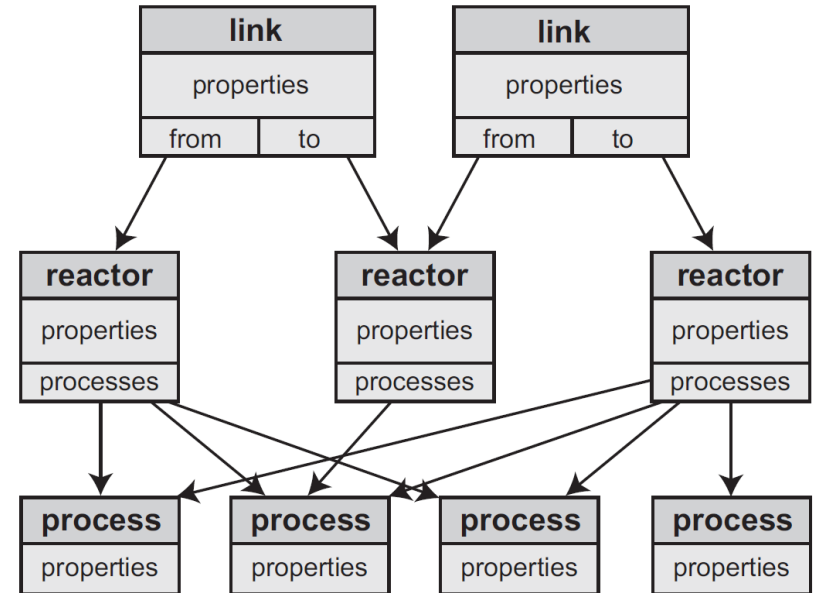
$$\frac{d}{dt} (A^k \mathbf{D}^k) = \boxed{A^k \mathbf{r}_{\mathbf{D}}^k} \quad \text{if}(Q, k_1, k_2) = \begin{cases} k_1 & \text{if } Q \geq 0 \\ k_2 & \text{if } Q < 0 \end{cases}$$

flow direct. logical:
notation to choose
concentration from
originating reactor

Multi-Reactor System

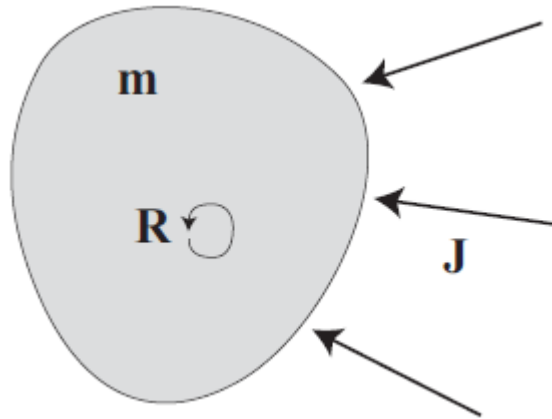
Implementation in Ecosim:

```
link <-  
  new(Class      = "link",  
       name      = "Metalimnion",  
       from      = "Epi",  
       to        = "Hypo",  
       flow      = 0,  
       qadv.gen   = 0,  
       qadv.spec  = list(C.POMD = expression(v.sed.POM*A),  
                         C.POMI = expression(v.sed.POM*A)),  
       qdiff.gen  = expression(A/h.meta*Kz),  
       qdiff.spec = 0)
```



Continuous System in Three Dimensions

chapter 3.4

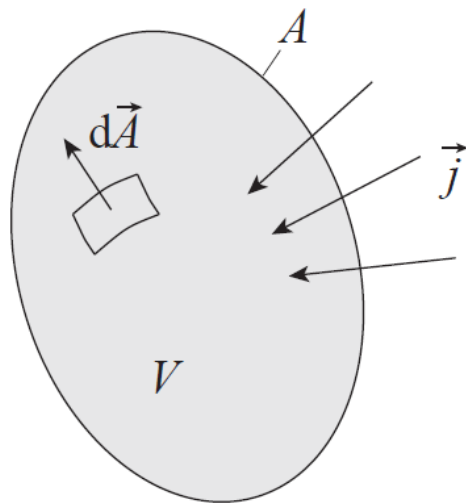


integral form:

$$\mathbf{m}(t_{\text{end}}) = \mathbf{m}(t_{\text{ini}}) + \int_{t_{\text{ini}}}^{t_{\text{end}}} \mathbf{J}(t) dt + \int_{t_{\text{ini}}}^{t_{\text{end}}} \mathbf{R}(t) dt$$

differential form:

$$\frac{d\mathbf{m}}{dt}(t) = \mathbf{J}(t) + \mathbf{R}(t)$$



$$\mathbf{m}(t) = \int_V \rho(\vec{x}, t) dV$$

$$\mathbf{J}(t) = - \int_A \vec{\mathbf{j}}(\vec{x}, t) \cdot d\vec{\mathbf{A}}$$

$$\mathbf{R}(t) = \int_V \mathbf{r}(\vec{x}, t) dV$$

total mass = volume
integral of mass density
(mass per volume)

net total flux =
surface integral
of flux density

net transformation rate =
volume integral of trans-
formation rate per volume

Continuous System in Three Dimensions

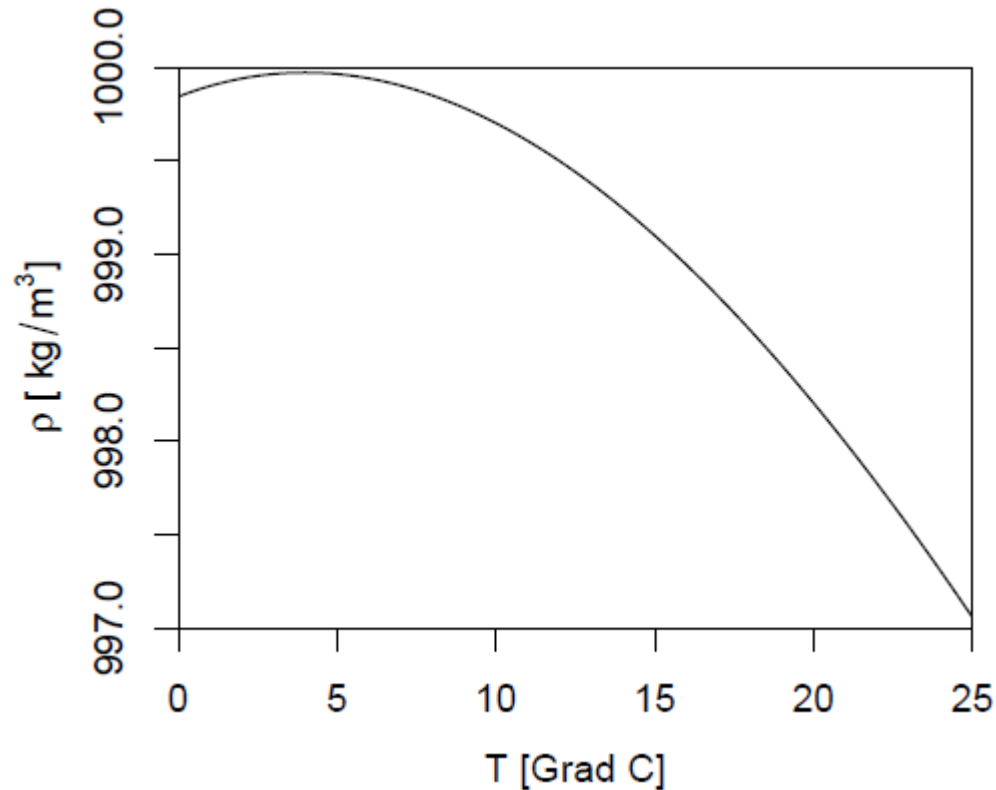
3D - Advection-Diffusion-Reaction equation

$$\begin{aligned} \frac{\partial C}{\partial t} + \overbrace{\frac{\partial(v_x C)}{\partial x} + \frac{\partial(v_y C)}{\partial y} + \frac{\partial(v_z C)}{\partial z}}^{\text{advection}} \\ = \underbrace{\frac{\partial}{\partial x} \left(D_x \frac{\partial C}{\partial x} \right) + \frac{\partial}{\partial y} \left(D_y \frac{\partial C}{\partial y} \right) + \frac{\partial}{\partial z} \left(D_z \frac{\partial C}{\partial z} \right)}_{\text{diffusion}} + \underbrace{r}_{\text{reaction}} \end{aligned}$$

Density of Water

chapter 6.1.1

The density of water is temperature dependent with a maximum at 4°C



Despite the weakness of this dependence, it has a very high effect on aquatic ecosystems.

Density Stratification of Water in a Lake

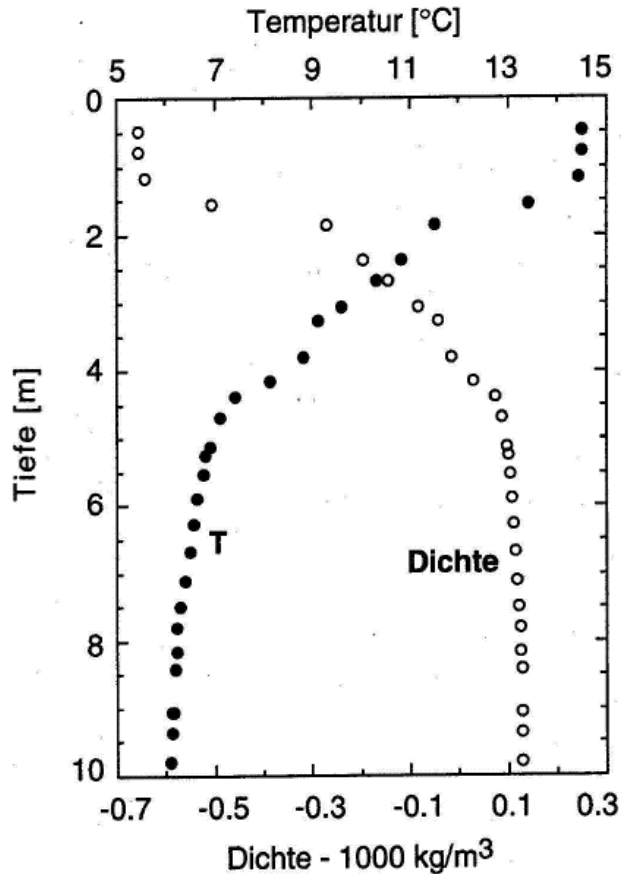
Spring/summer: Heating from the top leads to strong stratification of the water column:

- **Epilimnion:** The top 5-15 m of the lake contains warm water with stratification of the top layer during the day. However, wind and cooling during the night leads to a relatively well-mixed surface layer.
- **Hypolimnion:** The deep water zone in the lake is shielded from warming by the top layer.
- **Metalimnion:** In deep lakes, relatively thin zone with a very strong temperature gradient between epilimnion and metalimnion.

Autumn/winter: Cooling from the top and wind-induced turbulence often destroy stratification (partially or completely, depending on wind and depth).

With climate change this happens less often in larger Swiss lakes!
(e.g. Lake Zug)

Stratification of Lakes



Typical temperature and density profile
in spring (Lake Hallwil, spring 1992)

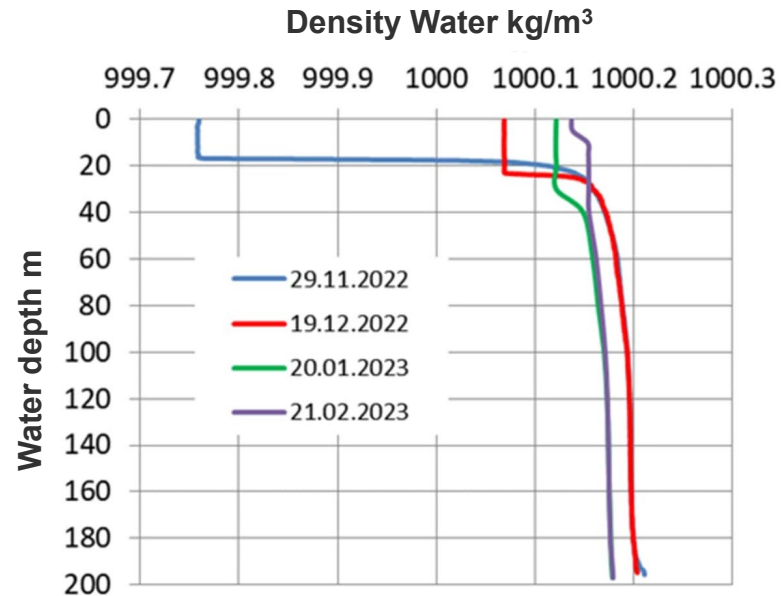
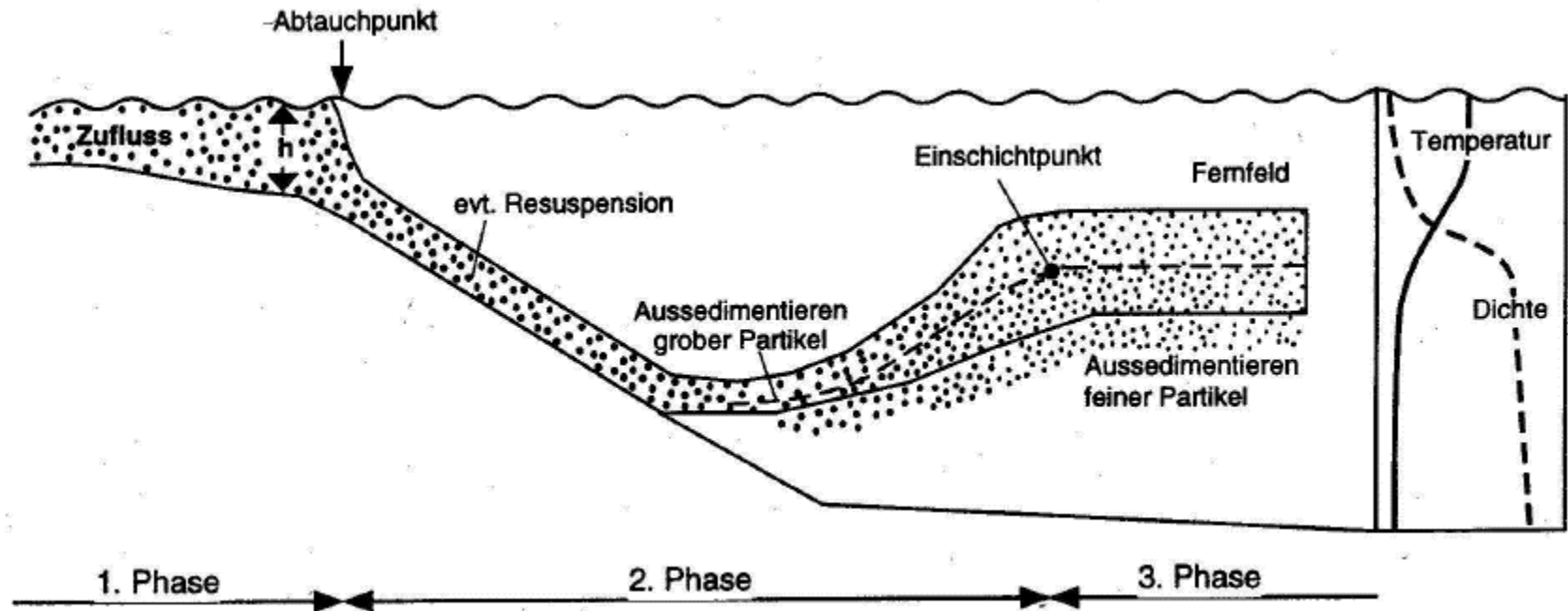


Abbildung 5: Dichteprofile des Zugersees in der Winterperiode 2022/23

https://zg.ch/dam/jcr:ba95d3d3-eabd-4694-8a71-0e9a85475748/20230821_TB%20Vorprojekt_Zirkulationunterst%C3%BCtzung_v2.1_ohne%20Anhang.pdf

Plunging of Inflows

Illustration of plunging process of inflow into a stratified water body (Wüest):



Plunging of Inflows

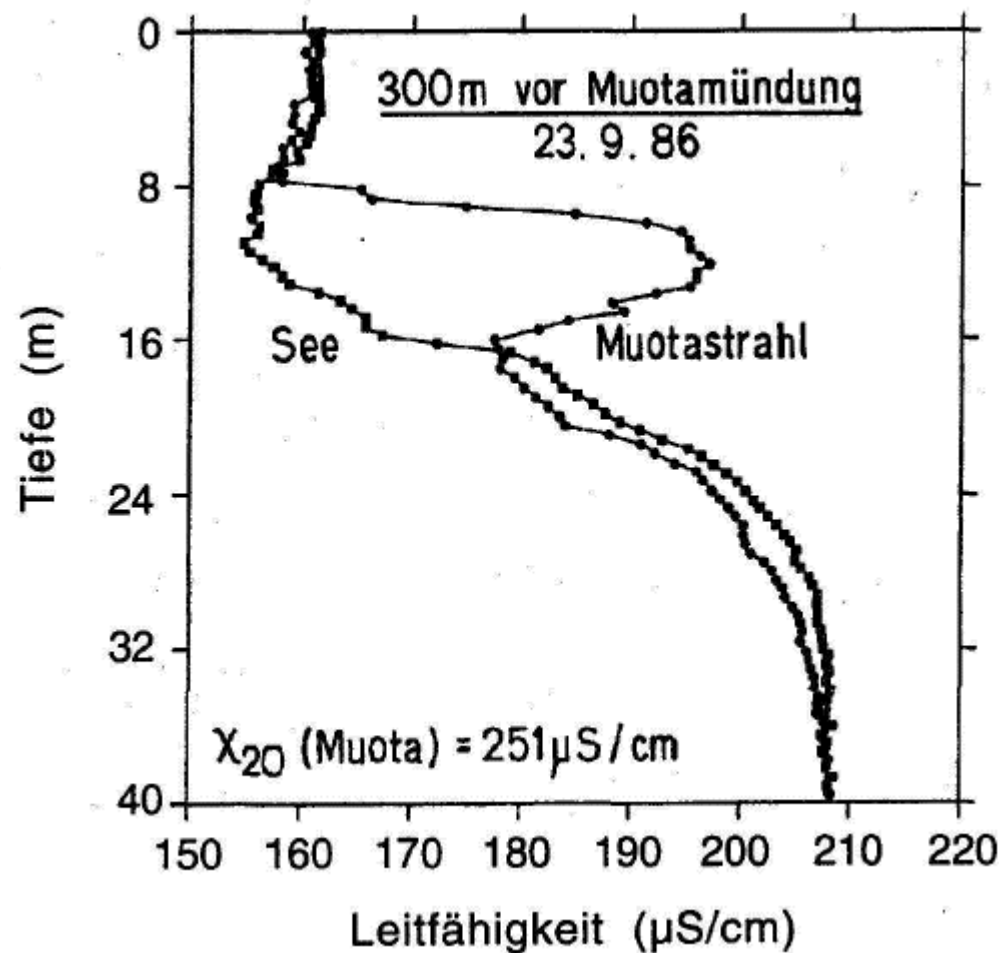
Another illustration of a plunging inflow (Seez into Walensee near Walenstadt):



Don't go swimming close to the inflow into a lake!

Plunging of Inflows

Conductivity signal of a plunging inflow (of Muota into lake Lucerne):



Horizontal Mixing

The diffusion part of the advection-diffusion-reaction equation reduced to two dimensions provides an approximate description of horizontal mixing:

$$\frac{\partial C}{\partial t} = K_{xy} \left(\frac{\partial^2 C}{\partial x^2} + \frac{\partial^2 C}{\partial y^2} \right)$$

↖
coefficient of horizontal turbulent diffusion

The solution of this equation for a substance point source of mass m , distributed over a depth h at time $t = 0$ is given by a normal distribution with linearly increasing variance in time:

$$C(x, y, t) = \frac{m}{h} \frac{1}{2\pi} \frac{1}{\sigma_{xy}(t)^2} \exp \left(-\frac{(x - x_0)^2 + (y - y_0)^2}{2\sigma_{xy}(t)^2} \right)$$

$$\sigma_{xy}(t) = \sqrt{2K_{xy}t}$$

Horizontal Mixing

The width of the substance distribution can roughly be estimated by 4 times the standard deviation and thus also increases with the square root of time:

$$L_{xy}(t) \approx 4\sigma_{xy}(t) = 4\sqrt{2K_{xy}t}$$

The maximum concentrations decreases with 1/time:

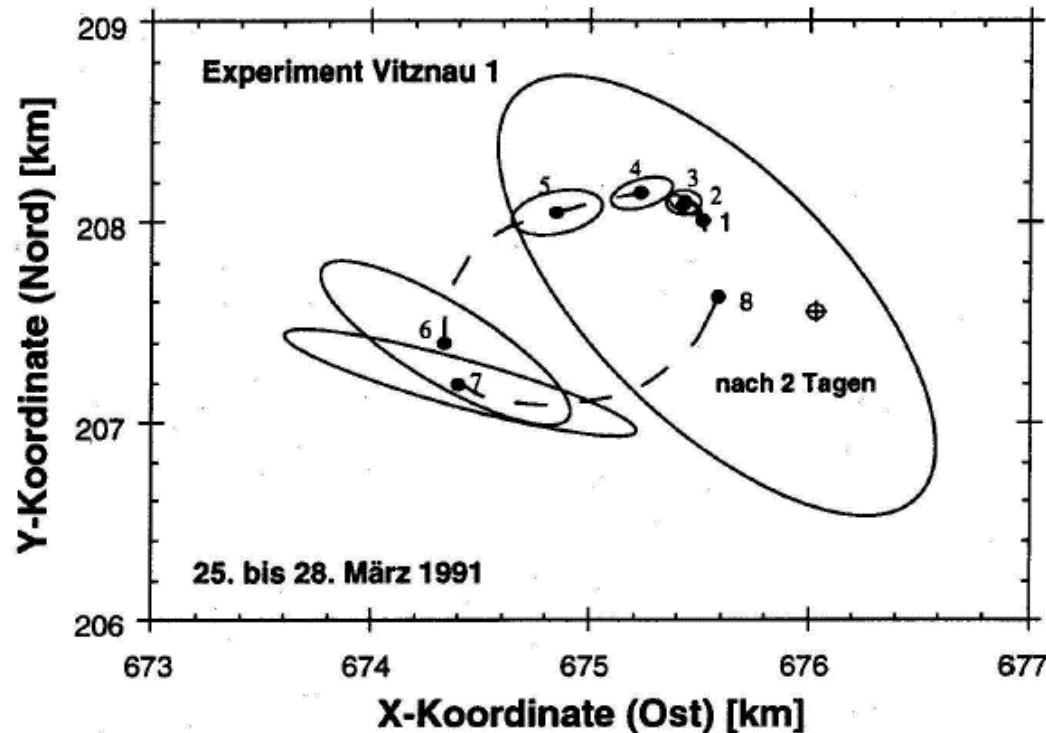
$$C_{\max} \approx \frac{m}{h} \frac{1}{4\sigma_{xy}(t)^2} = \frac{m}{h} \frac{1}{8K_{xy}t}$$

A typical range for the coefficient of horizontal turbulent diffusion is:

$$K_{xy} \approx 10^4 - 10^6 \text{ m}^2/\text{d}$$

Horizontal Mixing

An example of horizontal transport and mixing of an artificial tracer cloud:



Note that large eddies initially move the whole tracer cloud while when it becomes larger, they contribute to enlarging it. This leads to an increase of the apparent coefficient of turbulent diffusion with the size of the tracer cloud.

Vertical Mixing

The diffusion part of the advection-diffusion-reaction equation reduced to the vertical dimension provides an approximate description of vertical mixing:

$$\frac{\partial C}{\partial t} = K_z \frac{\partial^2 C}{\partial z^2}$$

 coefficient of vertical turbulent diffusion

The solution of this equation for a substance point source of mass m , distributed over the cross-sectional area A at time $t = 0$ is given by a normal distribution with linearly increasing variance in time:

$$C(z, t) = \frac{m}{A} \frac{1}{\sqrt{2\pi}} \frac{1}{\sigma_z(t)} \exp \left(-\frac{(z - z_0)^2}{2\sigma_z(t)^2} \right)$$

$$\sigma_z(t) = \sqrt{2K_z t}$$

Vertical Mixing

The overall size of the distribution can roughly be estimated by 4 times the standard deviation and thus also increases with the square root of time:

$$L_z(t) \approx 4\sigma_z(t) = 4\sqrt{2K_z t}$$

The maximum concentrations decreases with 1/time:

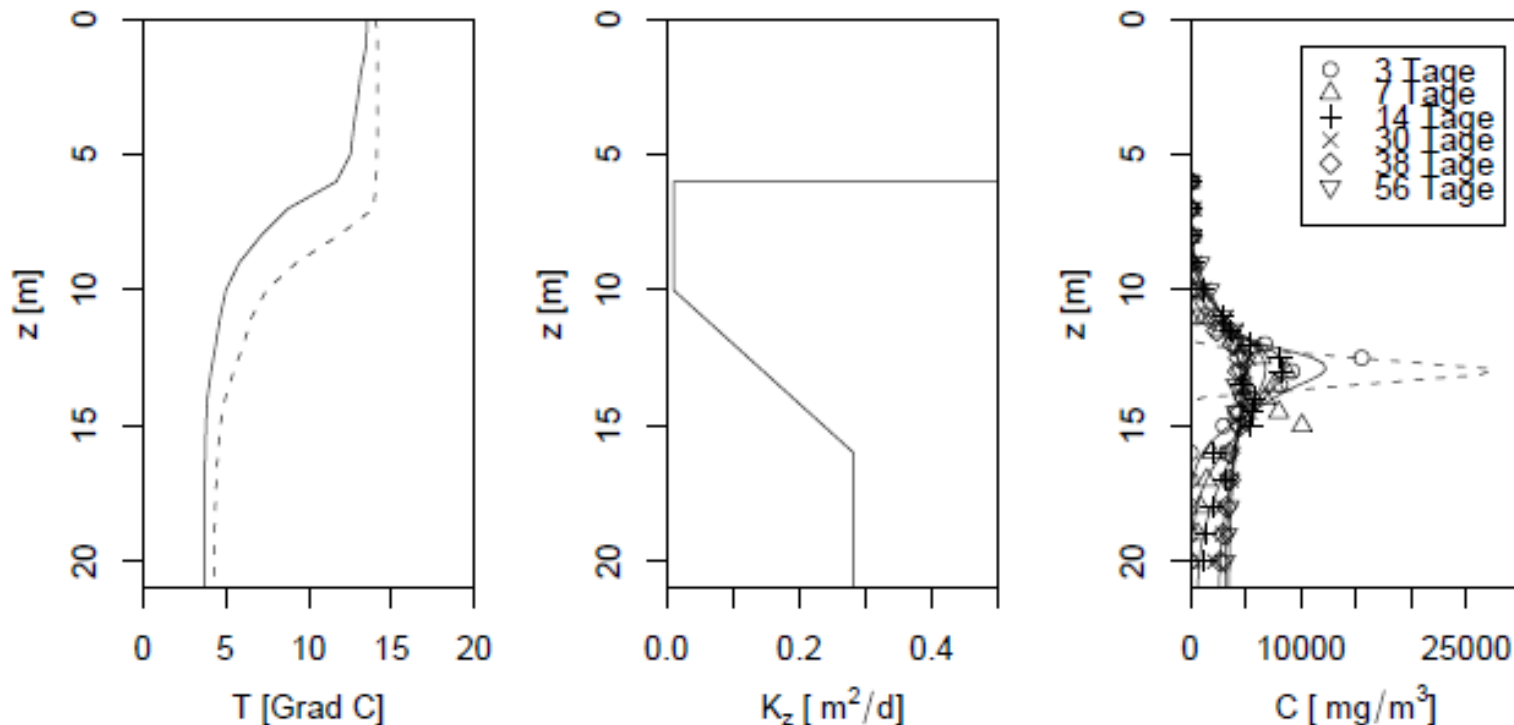
$$C_{\max} \approx \frac{m}{A} \frac{1}{2\sigma_z(t)} = \frac{m}{A} \frac{1}{2\sqrt{2K_z t}}$$

Typical ranges for the coefficient of vertical turbulent diffusion are:

$$\begin{array}{ll} K_{z,\text{epi}} \approx 100 \text{ m}^2/\text{d} & \approx \text{complete mixing} \\ K_{z,\text{meta}} \approx 0.01 - 0.1 \text{ m}^2/\text{d} & \left. \vphantom{\begin{array}{l} K_{z,\text{epi}} \\ K_{z,\text{meta}} \end{array}} \right\} \text{during stratification} \\ K_{z,\text{hypo}} \approx 0.1 - 10 \text{ m}^2/\text{d} & \end{array}$$

Vertical Mixing

The figure below demonstrates vertical diffusion of a tracer in lake Cadagno. The strong density gradient induced by the temperature gradient in the metalimnion prevents spreading of the tracer into the epilimnion:



Sedimentation

Imagine you are a particle floating around in a lake.

What do you think are important properties that influence if you are sedimented or you stay in suspension?

Sedimentation

chapter 6.2

Estimate of the sedimentation velocity: g : gravitational acceleration (LT^{-2}).

ρ_p : density of the particle (ML^{-3}).

ρ_w : density of water (ML^{-3}).

μ : dynamic viscosity of water ($ML^{-1}T^{-1}$).

d_p : particle diameter (L).

f_{st} : shape factor (1.0 for spheres).

$$v_{sed} = f_{st} \frac{g}{18} \frac{\rho_p - \rho_w}{\mu} d_p^2$$

Typical sedimentation velocities of algae, organic particles, clay and silt particles are in the following range:

$$v_{sed,ALG} \approx 0.0 - 0.3 \text{ m/d}$$

$$v_{sed,POM} \approx 0.2 - 2.5 \text{ m/d}$$

$$v_{sed,clay} \approx 0.3 - 1.0 \text{ m/d}$$

$$v_{sed,silt} \approx 3.0 - 30 \text{ m/d}$$

These estimate allow us to get an estimate of the sedimentation flux in the lake:

$$J_{sed,POM} = A_{sed} v_{sed,POM} C_{POM}$$

Gas Exchange

chapter 6.3

Equilibrium concentration according to Henry's law:

$$C_{w,eq}(C_a) = \frac{C_a}{H}$$

$C_{w,eq}(C_a)$ equilibrium concentration in water depending on T and p
 C_a concentration in the air
 H Henry's law constant, O₂ in water at 25°C, 1 atm: 0.032

If gas is produced or consumed:

Gas exchange flux between the atmosphere and the water

$$J_{\text{gasex}} = A v_{\text{ex}} \left(C_{w,eq}(C_a) - C_w \right)$$

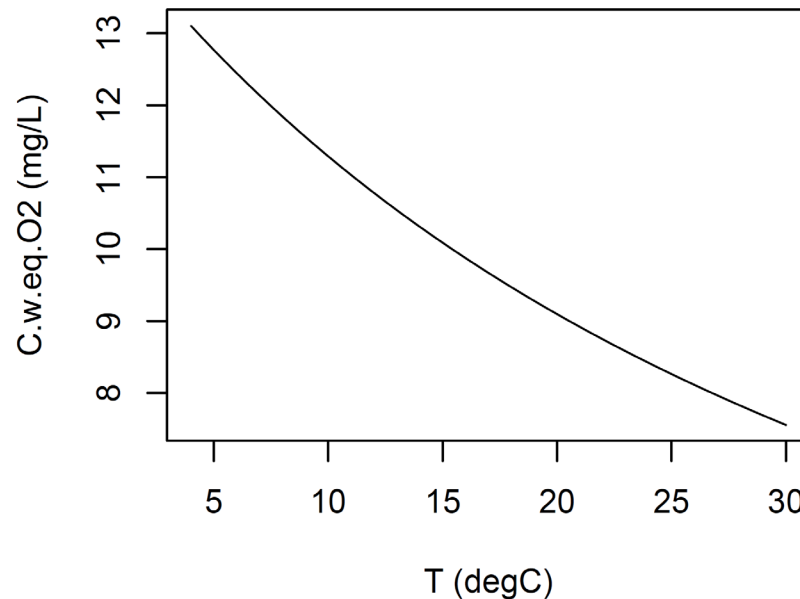
[M/T] [L²] [L/T] [M/L³]

Gas Exchange

chapter 6.3

Equilibrium concentration of O_2 in water depending on T and air pressure p :

$$C_{w,eq,O_2}(T, p) = \exp(7.7117 - 1.31403 \log(T/1^\circ\text{C} + 45.93)) \cdot 1\text{gO}/\text{m}^3 \cdot \frac{p}{1013.25 \text{ hPa}}$$



heat stress
can also be
due to lack of
oxygen

Two Box Lake Model

chapter 11.4

Important substance fluxes not related to directed water transport:

Exchange of oxygen at the lake surface between atmosphere and epilimnion:

$$J_{\text{gasex}} = A v_{\text{ex}} \left(C_{\text{w,eq}}(C_{\text{a}}) - C_{\text{w}} \right)$$

Exchange between epilimnion and hypolimnion through turbulent diffusion (all dissolved and suspended substances or organisms):

$$J^{\text{epi hypo}} = A_{\text{meta}} \frac{K_{\text{z,meta}}}{h_{\text{meta}}} (C^{\text{epi}} - C^{\text{hypo}})$$

Sedimentation flux of dead organic particles from epilimnion to hypolimnion:

$$J_{\text{POM}}^{\text{epi hypo}} = A_{\text{meta}} v_{\text{sed,POM}} C_{\text{POM}}$$

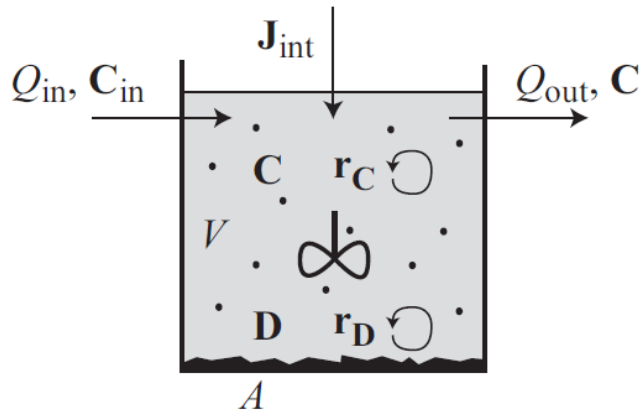
Note: For technical reasons (sediment layer part of box), we describe sedimentation from the hypolimnion to the sediment formally as a transformation process

Comment on Dissolved/Suspended and Attached Substances

In our mixed reactors, we distinguish substances that are dissolved or suspended in the water and transported with water flow and substances that are attached/sedimented/sessile on surfaces.

Substances in the water column are characterized by concentrations, C , (mass per volume), substances on surfaces by surface densities, D , (mass per area).

For processes involving substances of both types (e.g. mineralization of sedimented particles consuming dissolved oxygen from the water column), we have to consider this “geometric” difference.



We have to calculate total transformation rates $A r_D$, $V r_C$ to calculate fluxes between different types of substances.

This leads to conversion factors of A/V or V/A if converting per area into per volume rates or vice versa

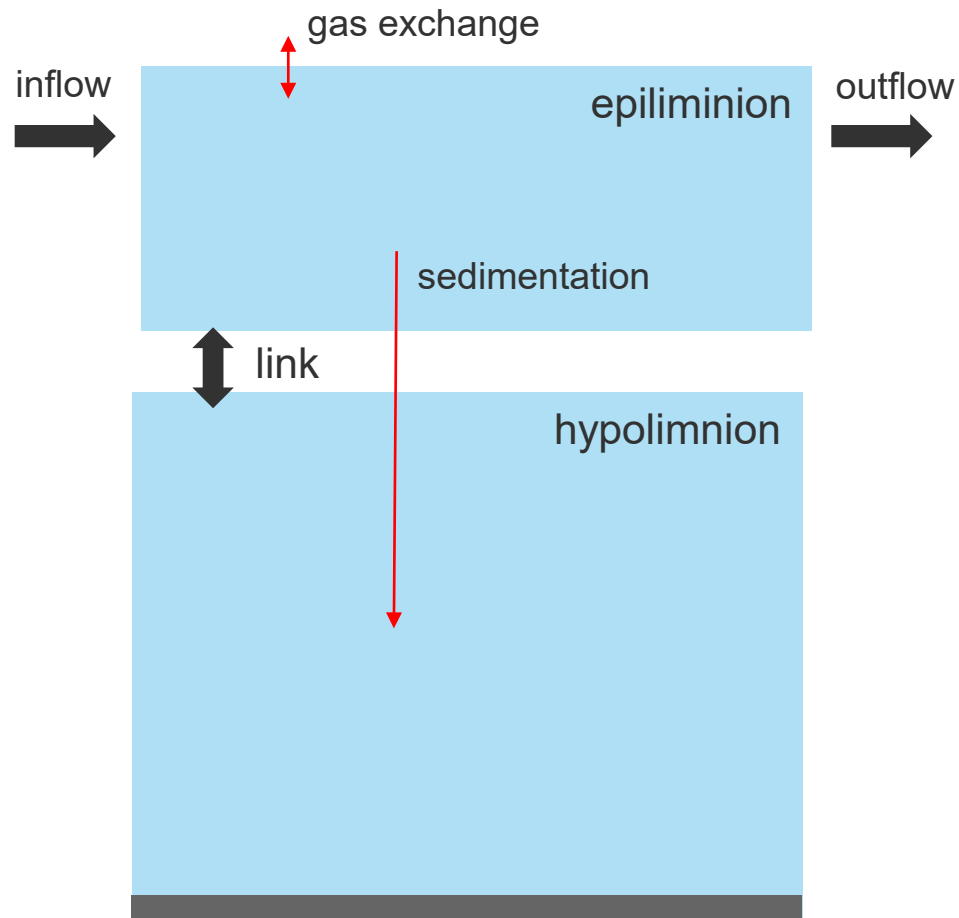
`ecosim` does this automatically for us!

State variables:

ALG, ZOO,

HPO_4^{2-} , NH_4^+ , NO_3^- , O_2

POM: degradable/inert,
suspended/sedimented



Processes and stoichiometry of the biogeochemical-ecological lake model

Process	Substances / Organisms									
	HPO_4^{2-} gP	NH_4^+ gN	NO_3^- gN	O_2 gO	ALG gDM	ZOO gDM	POMD gDM	POMI gDM	SPOMD gDM	SPOMI gDM
Growth of algae NO_3^-	–		–	+	1					
Growth of algae NH_4^+	–	–		+	1					
Respiration of algae	+	+		–	–1					
Death of algae	0/+	0/+		0/+	–1		$(1 - f_I)Y_{\text{ALG,death}}$	$f_I Y_{\text{ALG,death}}$		
Growth of zooplankton	+	+		–	$\frac{-1}{Y_{\text{ZOO}}}$	1	$\frac{(1 - f_I)f_e}{Y_{\text{ZOO}}}$	$\frac{f_I f_e}{Y_{\text{ZOO}}}$		
Respiration of zoopl.	+	+		–		–1				
Death of zooplankton	0/+	0/+		0/+		–1	$(1 - f_I)Y_{\text{ZOO,death}}$	$f_I Y_{\text{ZOO,death}}$		
Nitrification		–1	+	–						
Oxic mineral. of org. part.	+	+		–			–1			
Ox. min. of org. part. in sed.	+	+		–					–1	
Anox. min. of org. part. in sed.	+	+	–						–1	
Sed. of deg. org. part.							–1		1	
Sed. of inert org. part.								–1		1

$$\frac{dy}{dt} = g(y, t)$$

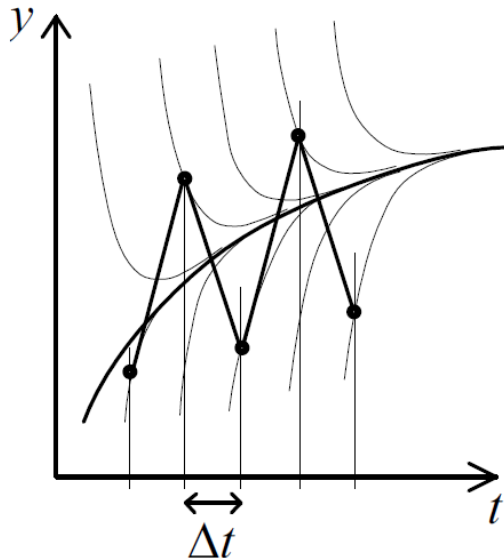
$$t_j = t_0 + j \cdot \Delta t \quad y_j = y(t_j)$$

explicit Euler scheme

$$y \rightarrow y_j$$

$$\frac{dy}{dt} \rightarrow \frac{y_{j+1} - y_j}{\Delta t}$$

$$y_{j+1} = y_j + \Delta t g(y_j, t_j)$$

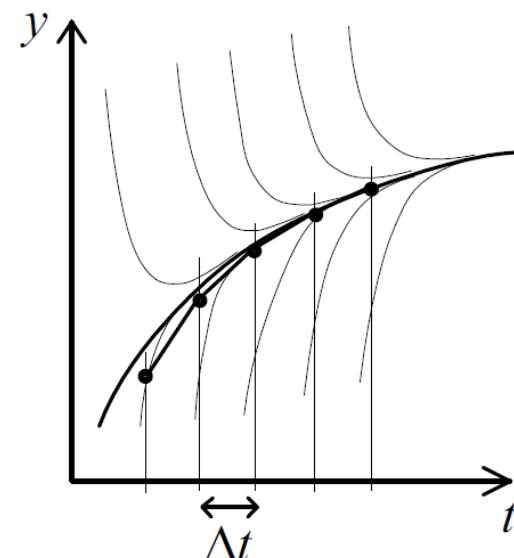


implicit Euler scheme

$$y \rightarrow y_j$$

$$\frac{dy}{dt} \rightarrow \frac{y_j - y_{j-1}}{\Delta t}$$

$$y_{j+1} = y_j + \Delta t g(y_{j+1}, t_{j+1})$$



→ explicit schemes are faster for a single step, implicit schemes can take larger steps.

Numerical implementation in `ecosim`

In `ecosim`, we use the R package `deSolve` that builds an interface to a large number of very sophisticated numerical ordinary differential equations solvers.

You can always get the description of R packages by entering the address

<https://cran.r-project.org/package=deSolve>

or

<https://cran.r-project.org/package=ecosim>

You find there the links to the reference manuals (you can also google for that)

Function `ode` in `deSolve` with a long list of numerical integrators available:

```
ode(y, times, func, parms,  
method = c("lsoda", "lsode", "lsodes", "lsodar", "vode", "daspk",  
           "euler", "rk4", "ode23", "ode45", "radau",  
           "bdf", "bdf_d", "adams", "impAdams", "impAdams_d", "iteration"), ...)
```

Function `calcre` in `ecosim` allows you to choose the method and passes further arguments to `ode` (...):

```
calcre(system, method="lsoda", ...)
```

There will be an **oral exam in the second week** after the semester
11./12. June 25 (4./5. June if needed)

It is your responsibility to register/deregister in time.

During the semester you will develop and implement your own model
(in teams of two people), interpret simulation results and perform a sensitivity
analysis. We will assign topics today.

Deadline for initial code submission: 08.05.25

Deadline for submission of R-files, results and interpretation: 23.05.25

This is **mandatory for being admitted** to the exam!

In the oral exam we will start with questions about your model before moving
on to other topics.

Please use the time in the exercises to ask questions and get help!
Don't do it last minute.

1. Introduction, principles of modelling environmental systems, mass balance in a mixed reactor, process table notation, simple lake plankton model
Exercise: R, ecosim-package, simple lake plankton model
Exercise: lake phytoplankton-zooplankton model
2. Process stoichiometry Exercises: analytical solution, calculation with stoichcalc
3. Biological processes in lakes
4. Physical processes in lakes, mass balance in multi-box and continuous systems Exercise: structured, biogeochemical-ecological lake model
Assignments: build your own model by implementing model extensions
5. Physical processes in rivers, bacterial growth, river model for benthic populations Exercise: river model for benthic populations, nutrients and oxygen
6. Stochasticity, uncertainty, Parameter estimation
Exercise: uncertainty, stochasticity
7. Existing models and applications in research and practice, examples and case studies, preparation of the oral exam, feedback

- Read the sections 11.3 and 11.4 about the didactical models
- think about your questions (also regarding your model assignments)